

# Simulation of Geophysical Flows.

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# Outline

- 1 Introduction
- 2 A naive derivation of the 1d Shallow Water model
- 3 Scalar conservation laws
- 4 Systems of conservation laws
- 5 Numerical methods

## Edanya group

### Edanya: Main Goals

Development of robust, reliable and **low computational cost** numerical tools for the simulation of geophysical flows and the prediction of emergency situations such as river floodings or oil spills, tsunamis, debris avalanches ...

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### Ingredients

- Mathematical models: based in **Shallow-water equations**.
- Numerical methods: **High order finite volume schemes**.
- HySEA: High Performance Cloud Computing software to simulate geophysical flows.

# HySEA: interdisciplinary platform

- Models based on geophysical flows with applications in:
  - Physical Oceanography
  - Marine Geology and Ecology
  - Tsunami Research
  - Civil and Hydraulic Engineering

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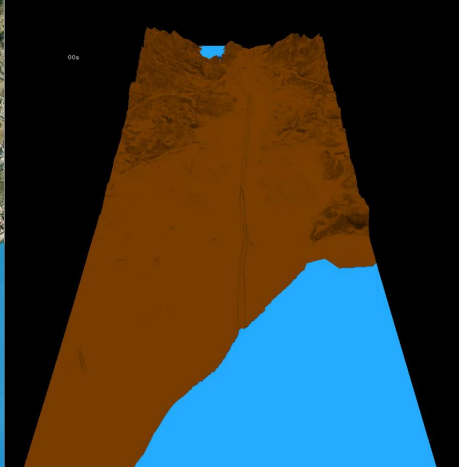
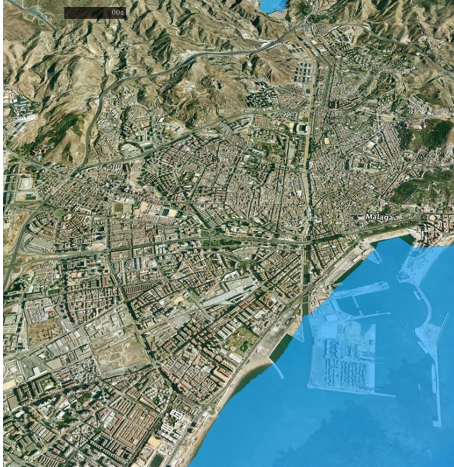
## Dambreak problem: Limonero Dam (close to Málaga (Spain))

- Resolution  $5 \text{ m} \times 5 \text{ m}$ .
- Number of cells: 1052224.
- Real simulated time: 20 min.
- Used scheme: Second order HLL or PVM-1U( $S_L, S_R$ ).
- Positivity of the water height is ensured.
- Graphics card: GeForce GTX 570. Speedup: 230 (1 Intel Core i7 920).

## Introduction

A naive derivation of the 1d Shallow Water model  
Scalar conservation laws  
Systems of conservation laws  
Numerical methods

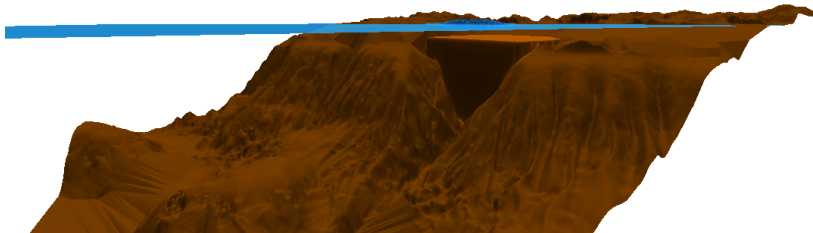
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# Tsunami generated by a submarine landslide in Sumatra

- Resolution  $20 \text{ m} \times 20 \text{ m}$ .
- Number of cells: 939500.
- Real simulated time: 12 min.
- Used scheme: Second order PVM-2U( $S_L, S_R, S_{int}$ ).

# Tsunami generated by a submarine landslide in Sumatra



## Shallow flow through a channel: notations

- Let us consider the free surface flow of a shallow layer of fluid through a channel with rectangular cross-section of constant width  $2A$ . Let  $L$  be the length.
- We consider a cartesian system of reference  $x, y, z$  such that the  $x$  axis is parallel the channel axis and  $z = 0$  corresponds to the undisturbed free surface level.
- Let  $H(x, y)$  be the depth of the channel measured from  $z = 0$ , i.e. the bottom is given by the surface  $\Gamma_b$  whose equation is:

$$z = -H(x, y) \quad 0 \leq x \leq L, \quad -A \leq y \leq A.$$

- Let us denote by  $\eta(x, y, t)$  the height of the free surface over the point  $(x, y, 0)$  at time  $t$ . The free surface at time  $t$  is given by the surface  $\Gamma_t^f$ , whose equation is:

$$z = \eta(x, y, t), \quad 0 \leq x \leq L, \quad -A \leq y \leq A.$$

- Let us denote by  $h(x, y, t)$  the thickness of the water layer over the point  $(x, y, 0)$  at time  $t$ , i.e.

$$h(x, y, t) = \eta(x, y, t) + H(x, y), \quad \forall x \in [0, L].$$

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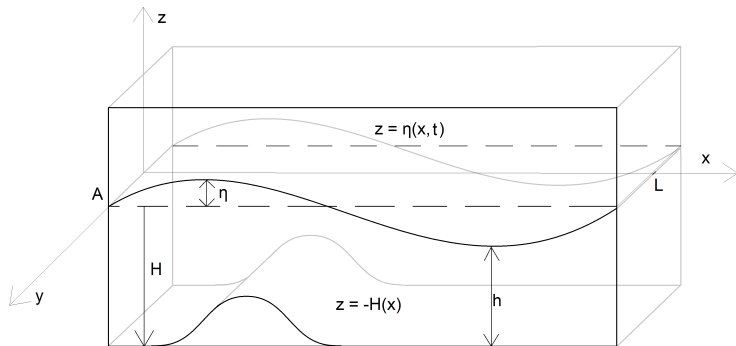
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- Let us denote by  $\Omega_t$  the 3D domain occupied by the water at time  $t$ :

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- Let us define:

$$\mathcal{Q} = \left\{ (x, y, z, t) \in \mathbb{R}^3 \times \mathbb{R}^+ \text{ s.t. } (x, y, z) \in \Omega_t \right\}.$$

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• **Fundamental Hypotheses** that are always assumed to derive the hyperbolic shallow water model.

• **Additional Hypotheses** that are used to derive the Saint Venant equations (shallow water model).

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## Fundamental Hypotheses

- **(FH1)** Water is assumed to be a **continuous medium**, i.e.
  - Water particles are identified with the dimensionless points of  $\Omega_t$ .
  - No mass can be assigned to the water particles, as there are uncountable infinitely many.
  - A mass  $m(\mathcal{O}, t)$  should be assigned to the water occupying any measurable subset  $\mathcal{O} \subset \Omega_t$ .
  - We assume the existence of a locally integrable function  $\rho : \mathcal{Q} \rightarrow \mathbb{R}^+$  such that

$$m(\mathcal{O}, t) = \int_{\mathcal{O}} \rho(x, y, z, t) dx dy dz,$$

for any measurable subset  $\mathcal{O}$  of  $\Omega_t$ . The function  $\rho$  is called the **density** function.

- We also assume the existence of a vector function  $\vec{u} : \mathcal{Q} \mapsto \mathbb{R}^3$  such that the velocity of the water particle which is at the point  $(x, y, z)$  at time  $t$  is given by:

$$\vec{u}(x, y, z, t) = (u_1(x, y, z, t), u_2(x, y, z, t), u_3(x, y, z, t)).$$

$\vec{u}$  is called the **velocity field**.

- Let us denote by

$$\vec{q}(x, y, z, t) := h(x, y, t)\vec{u}(x, y, z, t)$$

the **discharge**.

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## Fundamental Hypotheses

- (FH2) Density is assumed to be constant.

$$\rho(x, y, z, t) = \rho_0, \quad \forall (x, y, z, t) \in \mathcal{Q}.$$

- As a consequence:

$$m(\mathcal{O}, t) = \int_{\mathcal{O}} \rho(x, y, z, t) dx dy dz = \rho_0 |\mathcal{O}|,$$

where  $|\mathcal{O}|$  represents the Lebesgue measure of  $\mathcal{O}$ .

- (FH3) Water flow is assumed to be **incompressible**, i.e. if  $\mathcal{O}_t$  is the domain occupied at time  $t$  by the water particles that filled a subset  $\mathcal{O}_0$  at time  $t = 0$ , then

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- If it is assumed to be smooth enough, the incompressibility hypothesis is equivalent to assume that the divergence of  $\mathbf{v}$  is zero:

$$\nabla \cdot \mathbf{v} = \frac{\partial v_1}{\partial x} + \frac{\partial v_2}{\partial y} + \frac{\partial v_3}{\partial z} = 0.$$

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where  $|\mathcal{O}|$  represents the Lebesgue measure of  $\mathcal{O}$ .

- (FH3) Water flow is assumed to be **incompressible**, i.e. if  $\mathcal{O}_t$  is the domain occupied at time  $t$  by the water particles that filled a subset  $\mathcal{O}_0$  at time  $t = 0$ , then

$$|\mathcal{O}_t| = |\mathcal{O}_0|, \quad \forall t \in \mathbb{R}^+.$$

- If  $\vec{u}$  is assumed to be smooth enough, the incompressibility hypothesis is equivalent to assume that the divergence of  $\vec{u}$  is zero:

$$\nabla \cdot \vec{u} = \frac{\partial u_1}{\partial x} + \frac{\partial u_2}{\partial y} + \frac{\partial u_3}{\partial z} = 0.$$

## Fundamental Hypotheses

- (FH4) Pressure  $p$  is hydrostatic.

- The pressure of a fluid at rest increases linearly with depth. More precisely, it satisfies the equation:

$$\frac{\partial p}{\partial z} = -\rho g, \quad (1)$$

where  $g$  is the gravity acceleration.

- At the free surface, the pressure should be equal to the air pressure  $p_a$ :

$$p(x, y, \eta(x, y, t)) = p_a(x, y, \eta(x, y, t)). \quad (2)$$

- (1) and (2) constitute a Cauchy problem for a linear o.d.e for  $p(x, y, \cdot)$  whose solution is:

$$p(x, y, z, t) = p_a(x, y, \eta(x, y, t)) + \rho g (\eta(x, t) - z), \quad -H(x, y) \leq z \leq \eta(x, y, t).$$

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## Particular Hypotheses

- (PH1) The only external force exerted on the fluid is gravity and the only internal force is pressure. In particular, viscous and friction forces are neglected.
- (PH2) The depth function  $H$  only depends on the  $x$  variable:

$$H = H(x), \quad 0 \leq x \leq L.$$

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- (PH4) The free surface elevation  $\eta$  only depends on the variable  $x$  and  $t$ :

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- (PH5) The velocity field is such that:

There exists a velocity  $u(x, t)$  such that

$$v = u(x, t) \mathbf{e}_x, \quad w = 0.$$

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## Some calculations

- Given  $a$  and  $b$  such that  $0 \leq a < b \leq L$  and  $t > 0$ , let us define

$$\mathcal{O}_{a,b}^t = \{(x, y, z) \quad \text{s.t.} \quad a \leq x \leq b, \quad -A \leq y \leq A, \quad -H(x) \leq z \leq \eta(x, t)\}.$$

- The total mass of the water contained in  $\mathcal{O}_{a,b}^t$  is:

$$\begin{aligned} m_{a,b}^t &= \rho_0 \int_a^b \int_{-H(x)}^{\eta(x,t)} \int_{-A}^A 1 \, dy \, dz \, dx \\ &= 2A\rho_0 \int_a^b (\eta(x,t) + H(x)) \, dx \\ &= 2A\rho_0 \int_a^b h(x,t) \, dx. \end{aligned}$$

where

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- Notice that the pressure does not depend on  $y$ :

$$p = p_a + \rho g (\eta(x,t) - z) = \rho g (\eta(x,t) - z), \quad -H(x,y) \leq z \leq \eta(x,y,t).$$

- The outward normal unit vector at a point of the bottom  $\gamma = (x, y, -H(x))$  is:

$$\vec{n}(\gamma) = \frac{1}{\sqrt{1 + H'(x)^2}} (-H'(x), 0, -1).$$

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$$(x, y) \mapsto \gamma = (x, y, -H(x)) \in \Gamma_b,$$

some easy computation show that the surface element is given by:

$$d\gamma = \sqrt{1 + H'(x)^2} dx dy.$$

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$$\begin{aligned} \vec{P}'_{a,b} &= \int_{\Gamma_b} p(\gamma, t) \vec{n}(\gamma) d\gamma \\ &= 2A\rho_0 g \int_a^b (\eta(x, t) + H(x)) \frac{1}{\sqrt{1 + H'(x)^2}} (-H'(x), 0, -1) \sqrt{1 + H'(x)^2} dx \\ &= 2A\rho_0 g \left( - \int_a^b h(x, t) H'(x) dx, 0, - \int_a^b h(x, t) dx \right) \end{aligned}$$



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## Some calculations

- Given  $a \in [0, L]$ , the total pressure exerted on a cross-section

$$S_a^t = \{(a, y, z) \text{ s.t. } -A \leq y \leq A, \quad -H(a) \leq z \leq \eta(a, t)\}$$

by the fluid at its left is given by:

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- The total pressure exerted by the fluid at its right is:  $-\vec{P}_a^r$ .
- Exercise 1:** prove **Archimedes' principle**: the upward buoyant force exerted on a body immersed in a fluid at rest (i.e.  $\vec{u} = 0$  and  $\eta = 0$ ) is equal to the weight of the fluid that the body displaces.

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## Some calculations

- The mass of water per unit time passing through  $S_a^t$  from left to right is given by:

$$\begin{aligned}q_a^t &= \int_{S_a^t} \rho_0 \vec{u}(\gamma, t) \cdot (1, 0, 0) d\gamma \\&= 2A\rho_0 \int_{-H(a)}^{\eta(a,t)} u_1(a, t) dx \\&= 2A\rho_0 h(a, t) u_1(a, t) \\&= 2A\rho_0 q_1(a, t).\end{aligned}$$

- Obviously, the total mass of water per unit time passing through  $S_a^t$  from right to left is given by:

$$-q_a^t.$$

## Some calculations

- The mass of water per unit time passing through  $S_a^t$  from left to right is given by:

$$\begin{aligned}q_a^t &= \int_{S_a^t} \rho_0 \vec{u}(\gamma, t) \cdot (1, 0, 0) d\gamma \\&= 2A\rho_0 \int_{-H(a)}^{\eta(a,t)} u_1(a, t) dx \\&= 2A\rho_0 h(a, t) u_1(a, t) \\&= 2A\rho_0 q_1(a, t).\end{aligned}$$

- Obviously, the total mass of water per unit time passing through  $S_a^t$  from right to left is given by:

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## Some calculations

- More generally, if  $\varphi(x, t)$  represents the density function of any substance  $\mathcal{S}$  that is transported by the fluid, then the total amount of  $\mathcal{S}$  per unit time passing through  $S_a^t$  from right to left is given by:

$$\begin{aligned}
 Q_a^{S,t} &= \int_{S_a^t} \varphi(\gamma, t) \vec{u}(\gamma, t) \cdot (1, 0, 0) d\gamma \\
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- $Q_a^{S,t}$  is called the **flux** of  $\mathcal{S}$ .
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## Conservation Laws

- Mass conservation equation.

- Given  $a, b$  such that  $0 \leq a < b \leq L$  and  $t > 0$  the rate of change of the mass of water contained in  $\mathcal{O}_{a,b}^t$  is equal to the net flux of mass through the boundaries  $S_a$  and  $S_b$ .
- Using the notation introduced previously, the mathematical expression of this law is as follows:

$$\frac{d}{dt} m_{a,b}^t = q_a^t - q_b^t. \quad (3)$$

- And using the expressions that have been found:

$$2A\rho_0 \frac{d}{dt} \int_a^b h(x,t) dx = 2A\rho_0 q_1(a,t) - 2A\rho_0 q_1(b,t). \quad (4)$$

- Let us assume that  $h$  and  $q_1$  are smooth enough. Given  $t_0, t_1$  such that  $0 \leq t_0 < t_1$ , by integrating (4) between  $t_0$  and  $t_1$  and by using Barrow's rule, we obtain the **integral form of the mass conservation equation**:

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- On the other hand, we have:

$$\frac{d}{dt} \int_a^b h(x, t) dx = \int_a^b \frac{\partial h}{\partial t}(x, t) dx \quad (6)$$

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- **Momentum equation.**

- Newton's second law states that mass times acceleration is equal to the total force:

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or, equivalently,

$$\vec{F} = m\vec{a} = m\frac{d\vec{v}}{dt} = \frac{d}{dt}(m\vec{v}) = \frac{d\vec{q}}{dt},$$

where  $\vec{v}$  is the velocity and  $\vec{q}$ , the momentum.

- Coming back to the water flow through the channel, the momentum equation states that *the rate of change of the total momentum contained in  $\mathcal{O}_{a,b}^t$  is equal to the net flux of momentum through the boundaries  $S_a$  and  $S_b$  plus the total force.* Using the notation previously introduced:

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$$\int_a^b \int_{-A}^A \int_{-H(x)}^{\eta(x,t)} \rho_0(0, 0, -g) dz dy dx = (0, 0, -g m_{a,b}^t).$$

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- Using now the expressions of the different terms, we obtain:

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 2A\rho_0 \frac{d}{dt} & \left( \int_a^b h(x,t) u_1(x,t) dx, 0, \int_a^b \int_{-H(x)}^{\eta(x,t)} u_3(x,z,t) dz dx \right) = \\
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- The first component of the momentum equation is thus as follows:

$$\begin{aligned}
 \frac{d}{dt} \int_a^b q_1(x,t) dx & = q_1(a,t) u_1(a,t) - q_1(b,t) u_1(b,t) \\
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## Conservation Laws

- Taking into account that  $q_1 u_1 = q_1^2/h$  it can also be written as follows:

$$\begin{aligned} \frac{d}{dt} \int_a^b q_1(x, t) dx &= \frac{q_1^2}{h}(a, t) - \frac{q_1^2}{h}(b, t) \\ &+ \frac{g}{2}h(a, t)^2 - \frac{g}{2}h(b, t)^2 + g \int_a^b h(x, t)H'(x) dx. \end{aligned} \quad (10)$$

- By integrating from  $t = t_0$  to  $t_1$  and using Barrow's rule, we obtain the **integral form of the momentum equation**:

$$\begin{aligned} \int_a^b q_1(x, t_1) dx - \int_a^b q_1(x, t_0) dx &= \int_{t_0}^{t_1} \frac{q_1^2}{h}(a, t) dt - \int_{t_0}^{t_1} \frac{q_1^2}{h}(b, t) dt \\ \int_{t_0}^{t_1} \frac{g}{2}h(a, t)^2 dt - \int_{t_0}^{t_1} \frac{g}{2}h(b, t)^2 dt &+ g \int_{t_0}^{t_1} \int_a^b h(x, t)H'(x) dx dt. \end{aligned} \quad (11)$$

## Conservation Laws

- Taking into account that  $q_1 u_1 = q_1^2/h$  it can also be written as follows:

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## Conservation Laws

- If the functions are assumed to be smooth enough, using again Barrow's rule, we obtain from (9):

$$\int_a^b \left( \frac{\partial q_1}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q^2}{h} + \frac{g}{2} h^2 \right) - gh \frac{dH}{dx} \right) dx = 0. \quad (12)$$

- As  $a$  and  $b$  are arbitrary, we obtain the **differential form of the momentum equation**:

$$\frac{\partial q_1}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q^2}{h} + \frac{g}{2} h^2 \right) = gh \frac{dH}{dx}. \quad (13)$$

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## 1d Hyperbolic Shallow Water Model

- Putting together the differential forms of the mass and the momentum equation, we obtain the following PDE system with two equations and two unknowns ( $h$ ,  $q$ ):

$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial q}{\partial x} = 0. \\ \frac{\partial q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q^2}{h} + \frac{g}{2} h^2 \right) = gh \frac{dH}{dx}. \end{cases} \quad (14)$$

- This system has to be complemented with a set of initial conditions:

$$h(x, 0) = h_0(x), \quad q(x, 0) = q_0(x), \quad 0 \leq x \leq L,$$

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- If the bottom is flat, i.e.  $dH/dx = 0$ , the system can be written as follows:

$$\mathbf{u}_t + \mathbf{f}(\mathbf{u})_x = 0,$$

where

$$\mathbf{u} = \begin{pmatrix} h \\ q \end{pmatrix}, \quad \mathbf{f}(\mathbf{u}) = \begin{pmatrix} q \\ q^2/h + gh^2/2 \end{pmatrix}.$$

- This is a **system of conservation laws** and  $\mathbf{f}$  is its **flux function**.
- In the general case, it can be written as follows:

$$\mathbf{u}_t + \mathbf{f}(\mathbf{u})_x = \mathbf{S}(\mathbf{u}) \frac{dH}{dx},$$

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- **Exercise 2:** Check that the eigenvalues of the Jacobian of the flux function:

$$\frac{\partial \mathbf{f}}{\partial \mathbf{u}} = \begin{pmatrix} 0 & 1 \\ -\frac{q^2}{h^2} + gh & 2\frac{q}{h} \end{pmatrix} = \begin{pmatrix} 0 & 1 \\ -u^2 + gh & 2u \end{pmatrix}$$

are

$$\lambda_1 = u - \sqrt{gh}, \quad \lambda_2 = u + \sqrt{gh}.$$

- Notice that the eigenvalues are real and different, and thus the PDE system is strictly hyperbolic, if  $h > 0$ .

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- **Exercise 3:** Show that, for smooth solutions, the shallow water system is equivalent to:

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## Comments

- The shallow water model can be derived in a rigorous manner from the free surface incompressible Navier-Stokes through:
  - A dimensional analysis of the different terms to check those that can be neglected under the long waves assumption.
  - A vertical integration of the simplified 2d system.
- See, for instance: J.F. Gerbeau, B. Perthame, 2001.
- It is not necessary to assume that  $u_1 = u_1(x, t)$ . The only necessary assumption is that

$$|u_1(x, z, t) - \bar{u}_1(x, t)|$$

is small for every  $z$ , where:

$$\bar{u}_1(x, t) = \frac{1}{h(x, t)} \int_{-H(x)}^{\eta(x, t)} u_1(x, z, t) dz$$

is the depth-averaged velocity. In fact, the unknowns of the shallow water system are  $h$  and  $\bar{u}_1$ .

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$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial q}{\partial x} = 0. \\ \frac{\partial q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q^2}{h} + p(h) \right) = 0. \end{cases} \quad (16)$$

where

$$p = \frac{g}{2} h^2. \quad (17)$$

This is the 1d compressible Euler system, where  $h$  plays the role of the density and (17) is the law of state.

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and thus:

$$\begin{aligned} u_3(x, z, t) &= u_3(x, -H(x), t) + \int_{-H(x)}^{\eta(x,t)} \frac{\partial u_3}{\partial z} dz \\ &= u_3(x, -H(x), t) - \int_{-H(x)}^z \frac{\partial u}{\partial x} dz \\ &= u_3(x, -H(x), t) - (z + H(x)) \frac{\partial u}{\partial x}. \end{aligned}$$

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## Dimensionless form

- Let us consider some characteristic values  $x^*$ ,  $z^*$ ,  $t^*$ ,  $u^*$ , of the variables  $x$ ,  $z$ ,  $t$ ,  $u$ . For instance  $x^*$  is the length of the channel  $L$ ;  $z^*$ , the mean depth of the channel;  $u^*$ , the characteristic horizontal speed of the water; and

$$t^* = \frac{x^*}{u^*}.$$

- We consider the change of variables:

$$t = t^* t', \quad x = x^* x', \quad u = u^* u', \quad h = z^* h', \quad H = z^* H', \quad q = (z^* u^*) q'.$$

Some easy computations allow us to obtain the expression of the system in the dimensionless variables:

$$\begin{cases} \frac{z^*}{t^*} \frac{\partial h'}{\partial t'} + \frac{z^* u^*}{x^*} \frac{\partial q'}{\partial x'} = 0, \\ \frac{z^* u^*}{t^*} \frac{\partial q'}{\partial t'} + \frac{(z^* u^*)^2}{z^* x^*} \frac{\partial}{\partial x'} \left( \frac{q'^2}{h} \right) + \frac{(z^*)^2}{x^*} g \frac{\partial}{\partial x'} \left( \frac{h'^2}{2} \right) = \frac{(z^*)^2}{x^*} g h' \frac{dH'}{dx}. \end{cases}$$



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where

$$Fr = \frac{\sqrt{gz^*}}{u^*}$$

is the so-called **Froude number**.

- The basic hypothesis to derive the shallow water system from the incompressible Euler or Navier-Stokes system is

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Why this hypothesis has not been (apparently) necessary in the naive derivation?

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- Let us come back to the forgotten third component of the momentum equation:

$$2A\rho_0 \frac{d}{dt} \left( \int_a^b \int_{-H(x)}^{\eta(x,t)} u_3(x, z, t) dz dx \right) =$$

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- Using the equality

$$m'_{a,b} = \rho_0 g \int_a^b h(x, t) dx,$$

the expression found for  $u_3$ , Barrow's rule and the fact that  $a$  and  $b$  are arbitray, we end up with the equation:

$$\frac{\partial}{\partial t} \left( \frac{h^2}{2} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial x} \left( uhH \frac{\partial u}{\partial x} \right) = 0.$$

- But this equation cannot be derived from the shallow water system...

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- But this equation cannot be derived from the shallow water system...

## Dimensionless form

- Let us come back to the forgotten third component of the momentum equation:

$$2A\rho_0 \frac{d}{dt} \left( \int_a^b \int_{-H(x)}^{\eta(x,t)} u_3(x, z, t) dz dx \right) =$$

$$2A\rho_0 q(a, t)u_3(a, t) - 2A\rho_0 q(b, t)u_3(b, t) - 2Ag m_{a,b}^t + 2A\rho_0 g \int_a^b h(x, t) dx.$$

- Using the equality

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- Nevertheless, if we change to dimensionless variables, the left-hand side of this new equation becomes:

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## 2d shallow water model

- The general expression of the hyperbolic 2d shallow water model is as follows:

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_1}{\partial x} + \frac{\partial q_2}{\partial y} = 0, \\ \frac{\partial q_1}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q_1^2}{h} + gh^2 \right) + \frac{\partial}{\partial y} \left( \frac{q_1 q_2}{h} \right) = gh \frac{\partial H}{\partial x} + \tau_1, \\ \frac{\partial q_2}{\partial t} + \frac{\partial}{\partial x} \left( \frac{q_1 q_2}{h} \right) + \frac{\partial}{\partial y} \left( \frac{q_2^2}{h} + gh^2 \right) = gh \frac{\partial H}{\partial y} + \tau_2. \end{array} \right. \quad (19)$$

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## Friction forces

- The expression of the friction forces is given frequently by an empiric form. Usually it takes the form of a quadratic law:

$$-C(h)|\vec{u}|\vec{u}.$$

- Some possible expressions are the following:

• Chézy law:  $-\frac{g}{K^2}|\vec{u}|^2$ .

• Manning law:  $-\frac{g}{n^2}|\vec{u}|^2$ .

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- The friction with the air can also be taken into account using a quadratic expression of the wind velocity above the boundary layer.
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## Scalar conservation laws

- Let us consider a **Cauchy problem** of the form:

$$\begin{cases} u_t + f(u)_x = 0, & x \in \mathbb{R}, t \geq 0; \\ u(x, 0) = u_0(x), & x \in \mathbb{R}; \end{cases}$$

where:

- $f : \mathbb{R} \rightarrow \mathbb{R}$  is a smooth function;
  - $u_0 : \mathbb{R} \rightarrow \mathbb{R}$  is a known function.
- If  $u$  is a solution and the equation is integrated in a rectangle  $[a, b] \times [t_0, t_1]$  we obtain the **integral form** of the equation;

$$\int_a^b u(x, t_1) dx = \int_a^b u(x, t_0) dx + \int_{t_0}^{t_1} f(u(a, t)) dt - \int_{t_0}^{t_1} f(u(b, t)) dt.$$

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# Scalar conservation laws: examples

- **Burgers equation:**

$$u_t + \left( \frac{u^2}{2} \right)_x = 0.$$

- **Traffic models:**

$$u_t + (v \cdot u)_x = 0.$$

•  $u(x, t)$  traffic density;

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## Example: linear scalar conservation law

- Let us consider the particular case  $f(u) = au$ ,  $a \in \mathbb{R}$ :

$$\begin{cases} u_t + au_x = 0, & x \in \mathbb{R}, t \geq 0; \\ u(x, 0) = u_0(x), & x \in \mathbb{R}. \end{cases}$$

- If  $u$  is a solution, then  $u$  remains constant along every straight line of the family  $x = at + k$ ,  $k \in \mathbb{R}$ . In effect:

$$\frac{d}{dt}(u(at + k, t)) = a \frac{\partial u}{\partial x}(at + k, t) + \frac{\partial u}{\partial t}(at + k, t) = 0.$$

- These straight lines are the so-called **characteristic curves** of the PDE. Given an initial condition  $u_0$ , they allow us to easily compute the corresponding solution:

$$u(x, t) = u_0(x - at), \quad \forall x \in \mathbb{R}, t \geq 0.$$

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where  $a(u) = f'(u)$ .

- $a(u)$  is expected to play a similar role to  $a$  in the scalar case. This fact suggests to define the characteristic curves as the family of integral curves of the ODE:

$$\frac{dx}{dt} = a(u(x, t)).$$

- Notice that, unlike the linear case, the characteristic curves depend on the particular solution considered, as  $u$  explicitly appears in the expression of the ODE.

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- A solution remains constant along any characteristic curve:

$$\frac{d}{dt}(u(x(t), t)) = \left( \frac{\partial u}{\partial x} \frac{dx}{dt} + \frac{\partial u}{\partial t} \right) (x(t), t) = \left( \frac{\partial u}{\partial x} a(u) + \frac{\partial u}{\partial t} \right) (x(t), t) = 0.$$

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$$\begin{cases} \frac{dx}{dt} = a(u(x, t)); \\ x(0) = x_0. \end{cases} \quad (20)$$

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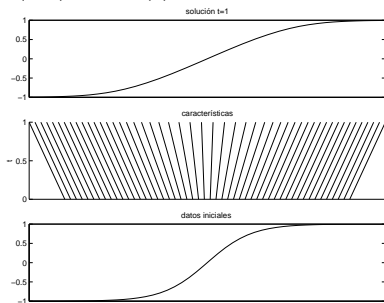
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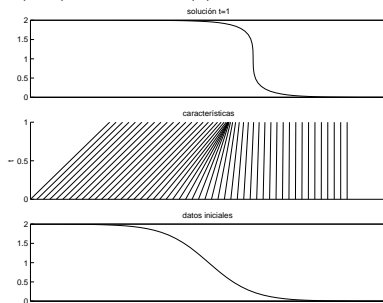
## Scalar conservation laws

- $u(x, 0) = \tanh(x)$ .



- The characteristic curves are divergent and the solution is smooth  $\forall t$ .

- $u(x, 0) = 1 - \tanh(x)$ .



- The characteristic are divergent and the solution becomes singular in a finite time.

### Weak solutions

## Scalar conservation laws

- **Goal:** to introduce a definition of weak solution allowing the solutions to become discontinuous.
- A function  $u : \mathbb{R} \times \mathbb{R}^+ \rightarrow \mathbb{R}$  is said to be **piecewise  $C^1$**  if there exists a finite number of smooth curves  $\Gamma_1, \dots, \Gamma_N$  outside which  $u$  is  $C^1$  and across which  $u$  has a jump discontinuity. Moreover, we assume that any line of discontinuity  $\Gamma_i$  has a parameterization of the form:

$$\Gamma_i = \{(\sigma_i(t), t), t \in I_i\},$$

where  $I_i$  is an interval of  $[0, \infty)$  and  $\sigma_i : I_i \subset \mathbb{R} \rightarrow \mathbb{R}$  a  $C^1(I_i)$  function. Let us denote by  $u^\pm$  the lateral limits of  $u$  at any point of a curve  $\Gamma_i$ :

$$\lim_{x \rightarrow \gamma^+} u(x, t_0) = u^+(\gamma, t_0), \quad \lim_{x \rightarrow \gamma^-} u(x, t_0) = u^-(\gamma, t_0).$$

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$$\forall (a, b) \subset \mathbb{R}, \forall (t_0, t_1) \subset \mathbb{R}^+.$$

- If  $u$  is a piecewise  $C^1$  function, the integrals in  $[a, b]$  appearing in the definition make sense.
- Nevertheless, there is a difficulty where the time integrals whenever there are a stationary discontinuity at  $x = a$  or  $x = b$ . This difficulty will be discussed later on.

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- **Example:** let us look for weak solutions of the form:

$$u(x, t) = \begin{cases} u_L & \text{if } x < \sigma(t); \\ u_R & \text{if } x > \sigma(t); \end{cases}$$

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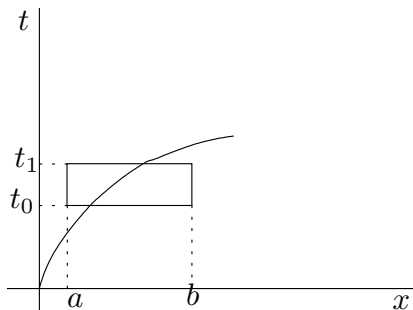
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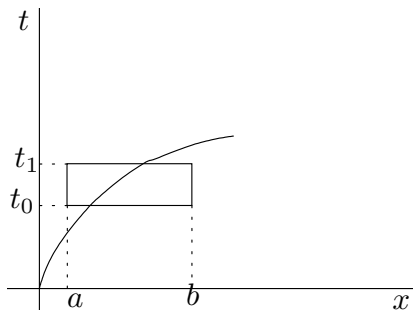


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Therefore,  $f(u)$  is continuous and thus the time integral

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appearing at Definition 1 is not ambiguous.

- It is possible to construct two conservation laws whose smooth solutions are the same but whose weak solutions are different. In other words, if a conservation law is reformulated using mathematical operations that are only valid for smooth solutions, it may happen that the corresponding Rankine-Hugoniot condition, and thus the weak solutions, are not equivalent.

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## Sel-similar solutions

- A function  $u$  is said to be a **self-similar solution** of the conservation law if it can be written in the form

$$u(x, t) = v\left(\frac{x - x_0}{t - t_0}\right)$$

for some  $x_0 \in \mathbb{R}$ ,  $t_0 \leq 0$  and a continuous function  $v : \mathbb{R} \rightarrow \mathbb{R}$ .

- If  $u$  is a self-similar solution, it is constant along the straight lines of the family:

$$\frac{x - x_0}{t - t_0} = C.$$

- For simplicity, let us take  $x_0 = t_0 = 0$ :

$$u(x, t) = v\left(\frac{x}{t}\right).$$

- Notice that a discontinuous weak solution of the form:

$$u(x, t) = \begin{cases} u_L & \text{if } x < st; \\ u_R & \text{if } x > st; \end{cases}$$

with  $s = [f(u)]/[u]$  is self-similar.

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- Let us consider Burgers equation with the initial condition:

$$u_0(x) = \begin{cases} -1 & \text{if } x < 0; \\ 1 & \text{if } x > 0; \end{cases}$$

- A solution of this Cauchy problem is the rarefaction wave.
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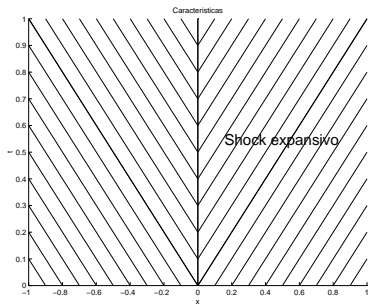
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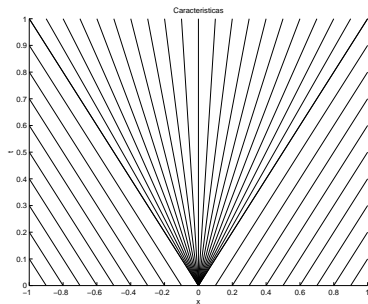
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$$u(x, t) = \text{sign}(x) = \begin{cases} -1 & x < 0 \\ 1 & x \geq 0 \end{cases} \quad v(x, t) = \begin{cases} -1 & x/t \leq -1 \\ x/t & -1 < x/t < 1 \\ 1 & x/t \geq 1 \end{cases}$$



Stationary discontinuity



Rarefaction wave

## Entropy condition

- It is necessary to have a criterion to decide whether a weak solution is physically meaningful or not.
- **Lax's entropy condition:** A discontinuity of a piecewise  $C^1$  solution  $u$  is said to be **admissible** if the Rankine-Hugoniot conditions are satisfied and

$$a(u^-) \geq s \geq a(u^+)$$

where  $u^-$  and  $u^+$  are the left and right limits at the discontinuity and  $s$ , the speed of propagation of the discontinuity.

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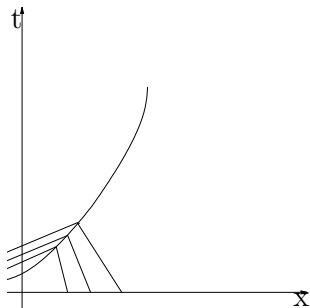
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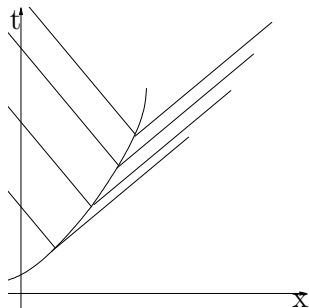
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Admissible discontinuity



Non-admissible discontinuity

# Entropy conditions

- **Vanishing viscosity method:**

- Let us consider a parabolic regularization of the conservation law:

$$\frac{\partial u_\varepsilon}{\partial t} + \frac{\partial}{\partial x} f(u_\varepsilon) = \varepsilon \frac{\partial^2 u_\varepsilon}{\partial x^2}.$$

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## Entropy conditions

- **Vanishing viscosity method:**

- Let us consider a parabolic regularization of the conservation law:

$$\frac{\partial u_\varepsilon}{\partial t} + \frac{\partial}{\partial x} f(u_\varepsilon) = \varepsilon \frac{\partial^2 u_\varepsilon}{\partial x^2}.$$

- The entropy solutions can be defined as those that can be obtained as a limit of a sequence  $u_\varepsilon$  of solutions of the parabolic regularization as  $\varepsilon$  goes to 0.
- **Definition:** a pair of functions  $(\eta, G)$  from  $\mathbb{R}$  to  $\mathbb{R}$  is said to be an **entropy pair** if:
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- It can be shown that the solutions of the parabolic regularizations satisfy the inequality

$$\eta(u_\varepsilon)_t + G(u_\varepsilon)_x \leq \varepsilon \frac{\partial^2}{\partial x^2} \eta(u_\varepsilon).$$

- Definition 2:** A weak solution is said to be an **entropy solution** if it satisfies the inequality:

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- It can be shown that a piecewise  $C^1$  solution is an entropy solution if and only if the inequality

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# Entropy condition

- For scalar conservation laws, there always exist infinitely many entropy pairs.
- If  $f$  is strictly convex or concave, the definition of entropy solution is independent of the chosen entropy pair.
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## The Riemann Problem

- Given a scalar conservation law, a **Riemann Problem** is a Cauchy problem whose initial condition has the form:

$$u(x, 0) = \begin{cases} u_L & x < 0 \\ u_R & x \geq 0 \end{cases}$$

- If  $f$  is strictly convex or concave, the solution of the Riemann problem is given by:
  - If  $u(u_L) < u(u_R)$  the solution is the rarefaction wave:

$$u(x, t) = \begin{cases} u_L & \text{if } x < a(u_L)t \\ a^{-1}(x/t) & \text{if } a(u_L)t < x < a(u_R)t \\ u_R & \text{if } x > a(u_R)t \end{cases}$$

where  $a(u) = f'(u)$  is the characteristic speed.

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$$s = f(u_R) - f(u_L) / u_R - u_L$$

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- Notice that the solution of the Riemann problem is self-similar in all the cases. In what follows it will be denoted by:

$$u(x, t) = V\left(\frac{x}{t}; u_L, u_R\right).$$

- As a consequence, observe that the solution of the Riemann problem at  $x = 0$  is constant.

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## The linear case

- Let us consider the PDE system:

$$\mathbf{u}_t + \mathcal{A}\mathbf{u}_x = 0,$$

where

$$\mathbf{u}(x, t) = \begin{bmatrix} u_1(x, t) \\ \vdots \\ u_N(x, t) \end{bmatrix},$$

and  $\mathcal{A}$  is a constant  $N \times N$  matrix.

- Let us assume that the system is strictly hyperbolic, i.e.  $\mathcal{A}$  has  $N$  different real eigenvalues  $\lambda_1 < \dots < \lambda_N$ . Let  $\{\mathbf{r}_1, \dots, \mathbf{r}_N\}$  be a basis of associated eigenvectors.



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- Let  $\mathcal{K}$  be the matrix whose columns are given by the eigenvectors  $\mathbf{r}_i$ . One has:

$$\mathcal{A} = \mathcal{K}\Lambda\mathcal{K}^{-1},$$

where  $\Lambda$  is the diagonal matrix whose coefficients are the eigenvalues.

- If the change of variables

$$\mathbf{v} = \mathcal{K} \cdot \mathbf{u},$$

is applied, the system reduces to  $N$  uncoupled scalar linear conservation laws:

$$\frac{\partial v_i}{\partial t} + \lambda_i \frac{\partial v_i}{\partial x} = 0, i = 1, \dots, N.$$

- The components of  $\mathbf{v}$  are called the **characteristic variables**, while those of  $\mathbf{u}$  are called the **conserved variables**.

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- To solve the Cauchy problem with initial condition

$$\mathbf{u}(x, 0) = \mathbf{u}_0(x),$$

- First, the expression of the initial condition in the eigenvector basis is computed:

$$\mathbf{u}_0(x) = \sum_{i=1}^N v_i^0(x) \mathbf{r}_i.$$

- Next, the uncoupled conservation laws are solved:

$$v_i(x, t) = v_i^0(x - \lambda_i t), \quad i = 1, \dots, N.$$

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# Systems

- In particular, the solution of the Riemann problem whose initial conditions are given by two states  $\mathbf{u}_L, \mathbf{u}_R$  consist of  $N$  discontinuities issuing from the origin and travelling at constant speeds  $\lambda_1, \dots, \lambda_N$ .
- These discontinuities link  $N + 1$  states:

$$\mathbf{u}_0 = \mathbf{u}_L, \mathbf{u}_1, \dots, \mathbf{u}_{N-1}, \mathbf{u}_N = \mathbf{u}_R.$$

- If

$$\mathbf{u}_L = \sum_{i=1}^N \alpha_i^L \mathbf{r}_i, \quad \mathbf{u}_R = \sum_{i=1}^N \alpha_i^R \mathbf{r}_i,$$

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- The Jacobian matrix of the flux function  $F$  is given by:

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## Eigenvalues, characteristic fields

- Let us assume that the system is strictly hyperbolic: for every state  $U$ ,  $\mathcal{A}(\mathbf{u})$  has  $N$  different real eigenvalues  $\lambda_1(\mathbf{u}) < \dots < \lambda_N(\mathbf{u})$ . Let  $\mathbf{r}_1(\mathbf{u}), \dots, \mathbf{r}_N(\mathbf{u})$  be associated eigenvectors.

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- If

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## Speed of propagation of small waves

- The eigenvalues are the speed of propagation of small waves: let us consider a sinusoidal perturbation of a constant state  $\mathbf{u}_0$ :

$$\mathbf{u}_0 + \bar{\mathbf{u}} \cos(\omega t),$$

which is small in the sense that  $\bar{\mathbf{u}} \ll \mathbf{u}_0$ .

- Let us look for a solution of the form:

$$\mathbf{u}(x, t) = \mathbf{u}_0 + \bar{\mathbf{u}} \cos(\omega(x - ct)),$$

i.e. the perturbation propagating at speed  $c$ .

- One has:

$$\begin{aligned} 0 &= \mathbf{u}_t + \mathbf{f}(\mathbf{u})_x \\ &= \mathbf{u}_t + \mathcal{A}(\mathbf{u})\mathbf{u}_x \\ &= (-c\bar{\mathbf{u}} + \mathcal{A}(\mathbf{u})\bar{\mathbf{u}}) \omega \sin(\omega(x - ct)) \\ &\cong (-c\bar{\mathbf{u}} + \mathcal{A}(\mathbf{u}_0)\bar{\mathbf{u}}) \omega \sin(\omega(x - ct)). \end{aligned}$$

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which is small in the sense that  $\bar{\mathbf{u}} \ll \mathbf{u}_0$ .

- Let us look for a solution of the form:

$$\mathbf{u}(x, t) = \mathbf{u}_0 + \bar{\mathbf{u}} \cos(\omega(x - ct)),$$

i.e. the perturbation propagating at speed  $c$ .

- One has:

$$\begin{aligned} 0 &= \mathbf{u}_t + \mathbf{f}(\mathbf{u})_x \\ &= \mathbf{u}_t + \mathcal{A}(\mathbf{u})\mathbf{u}_x \\ &= (-c\bar{\mathbf{u}} + \mathcal{A}(\mathbf{u})\bar{\mathbf{u}}) \omega \sin(\omega(x - ct)) \\ &\cong (-c\bar{\mathbf{u}} + \mathcal{A}(\mathbf{u}_0)\bar{\mathbf{u}}) \omega \sin(\omega(x - ct)). \end{aligned}$$

## Speed of propagation of small waves, jump condition

- We find an approximate solution if there exists  $i \in \{1, \dots, N\}$  such that

$$c = \lambda_i(\mathbf{u}_0)$$

and  $\bar{\mathbf{u}}$  is an associated eigenvector.

- The Rankine-Hugoniot condition writes as follows:

$$[\mathbf{f}(\mathbf{u})] = s[\mathbf{u}],$$

where  $s$  is the speed of propagation of the shock.

- Lax's entropy condition:** a shock is admissible if there exists  $i \in \{1, \dots, N\}$  such that:

$$\lambda_i(\mathbf{u}^-) > s > \lambda_{i-1}(\mathbf{u}^-), \quad \lambda_{i+1}(\mathbf{u}^+) > s > \lambda_i(\mathbf{u}^+)$$

if the  $i$ -th characteristic field is genuinely nonlinear, or

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## Characteristic curves, simple waves, Riemann problem

- Given a solution  $\mathbf{u}$ , the **characteristic curves** are the integral curves of the  $N$  ODE systems:

$$\frac{dx}{dt} = \lambda_i(\mathbf{u}(x(t), t)), \quad i = 1, \dots, N.$$

- Associated to every family of characteristic curves, and depending of their nature, one can construct solutions that are:
  - contact discontinuities that propagate following a characteristic curve, as in the linear scalar case, if the characteristic field is linearly degenerate;
- These solutions are called **simple waves**.
- If all the characteristic fields are either genuinely nonlinear or linearly degenerate, the solution of the Riemann problem consists of a  $N$  simple waves linking  $N + 1$  intermediate states.

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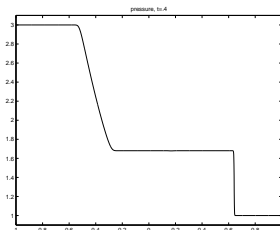
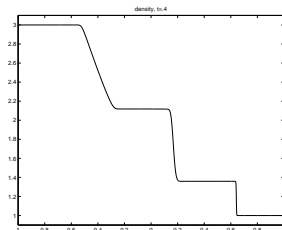
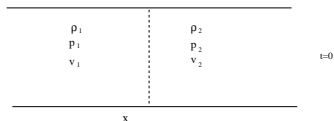
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## Example

Example: Euler system with initial conditions

$$\rho(x, 0) = \begin{cases} \rho_1 & x < 0 \\ \rho_2 & x > 0 \end{cases} \quad v(x, 0) = \begin{cases} v_1 & x < 0 \\ v_2 & x > 0 \end{cases} \quad p(x, 0) = \begin{cases} p_1 & x < 0 \\ p_2 & x > 0 \end{cases}$$



## The shallow water model

- In the particular case of the shallow water system:

$$\mathbf{u} = \begin{pmatrix} h \\ q \end{pmatrix}, \quad \mathbf{f}(\mathbf{U}) = \begin{pmatrix} q \\ q^2/h + gh^2/2 \end{pmatrix}.$$

$$\frac{\partial \mathbf{f}}{\partial \mathbf{u}} = \begin{pmatrix} 0 & 1 \\ -\frac{q^2}{h^2} + gh & 2\frac{q}{h} \end{pmatrix} = \begin{pmatrix} 0 & 1 \\ -u^2 + gh & 2u \end{pmatrix}$$

$$\lambda_1(\mathbf{u}) = u - \sqrt{gh}, \quad \lambda_2(\mathbf{u}) = u + \sqrt{gh}.$$

- The eigenvalues are real and different if  $h > 0$ . They are the speed of propagation of small waves.
- **Exercise 4:** Compute the characteristic fields and check that they are genuinely non-linear.
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# The shallow water model

- Let  $Fr$  be the Froude number:

$$Fr = \frac{|u|}{\sqrt{gh}}.$$

- If  $Fr > 1$ , then both eigenvalues  $u \pm \sqrt{gh}$  have the same sign of  $u$ : the small waves travel in the flow sense. The flow is said to be **supercritical** (is the equivalent concept to supersonic in the case of a compressible gas).
- If  $Fr < 1$  the signs of the eigenvalues are different: some small waves travel in the flow sense and some other in the other sense. The flow is said to be **subcritical**.
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$$\begin{aligned} s[h] &= [q], \\ s[q] &= \left[ \frac{q^2}{h} + g \frac{h^2}{2} \right] \end{aligned}$$

- The first equation can be rewritten as follows:

$$v^+ h^+ = v^- h^-,$$

where

$$v^+ = u^+ - s, \quad v^- = u^- - s$$

are the speed of the particles relative to the shock speed at both sides of the shock. Therefore, both relative speeds have the same sign.

- The fluid particle cross the shock from left to right if  $v^\pm > 0$  and from right to left if  $v^\pm < 0$ .
- It can be shown that the entropy condition is equivalent to:

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## Shocks in shallow waters



TMAA SCAN

### F-4 Phantom II Caught Breaking the Sound Barrier.

Using a 35mm camera, a telephoto lens and ASA 400 film, Pat Maloney, an engineering planner, photographed an F-4 Phantom II at the moment it broke the sound barrier at the Annual Point Magu Naval Air Station Air Show. "The photograph of the visible shock wave is rare," stated Maloney. "It required a humid day, split second timing and no small measure of luck." Maloney frequently practices photography at the many air shows he attends.



The Military Aircraft Archive:  
<http://www.milair.simplenet.com>



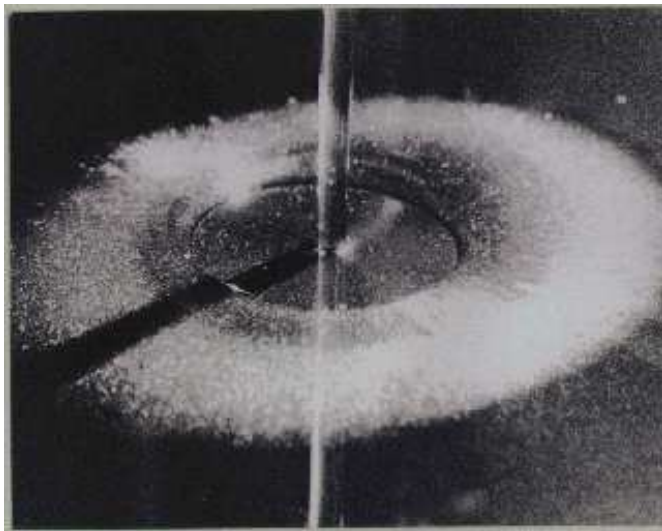
## Shocks in shallow waters



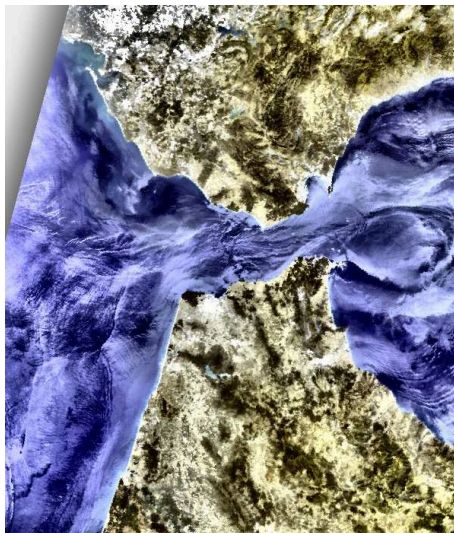
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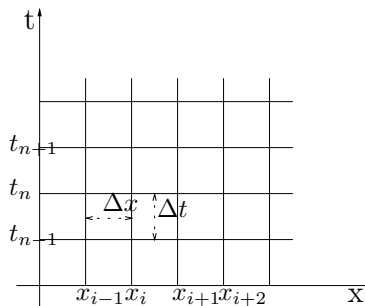
- Internal tides in the Strait of Gibraltar.

## Finite differences

- Let us consider the Cauchy problem:

$$\begin{cases} \frac{\partial u}{\partial t} + a \frac{\partial u}{\partial x} = 0, & x \in \mathbb{R}, t > 0; \\ u(x, 0) = u_0(x), & x \in \mathbb{R}. \end{cases}$$

- For simplicity, we consider a uniform mesh of the half-plane  $t \geq 0$ .

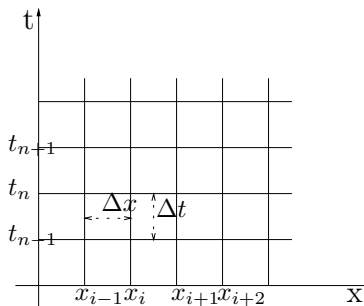


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## Finte differences

- The partial derivatives appearing at the equation are approached using only points of the mesh.
- Space partial derivatives:

$$\frac{\partial u}{\partial x}(x_i, t_n) = \frac{u(x_{i+1}, t_n) - u(x_i, t_n)}{\Delta x} + O(\Delta x),$$

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# Finite difference

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- Backward difference method:

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- Centered method:

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# Las's equivalence theorem

- **Theorem:** For linear problems, a **consistent** numerical method is **convergent** if and only if it is **stable**.
- **Consistency:** The exact values of a smooth solutions at the mesh points have to satisfy the finite difference equation with errors that converge to 0 as  $\Delta x$ ,  $\Delta t$  tend to 0. These errors are the so-called **local discretization errors**.

For instance, for the forward difference method:

$$L_i^* = \frac{u(x_i, t_{n+1}) - u(x_i, t_n)}{\Delta t} + a \frac{u(x_{i+1}, t_n) - u(x_i, t_n)}{\Delta x} = O(\Delta x + \Delta t).$$

For a second order method (e.g. Lax-Wendroff):

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# Stability

- **Stability:** The local discretization errors and the round-off errors are not dramatically amplified as the numerical method advances in time.
  - As the exact solution satisfies:

$$\sup\{|u(x, t)|, x \in \mathbb{R}\} = \sup\{|u_0(x)|, x \in \mathbb{R}\}, \quad \forall t > 0,$$

a reasonable stability criterion is the following: given  $T > 0$ , there exists  $K > 0$  independent of  $\Delta t$  and  $\Delta x$  such that:

$$\sup_{i \in \mathbb{Z}} |u_i^n| \leq K \sup_{i \in \mathbb{Z}} |u_i^0|, \quad \forall n \leq \frac{T}{\Delta t}.$$

- Let  $u^n$  be the piecewise constant function taking the value  $u_i^n$  in  $[x_i - \Delta x/2, x_i + \Delta x/2)$ . Using this function, the previous inequality writes as follows:

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- It is possible to give an alternative definition of stability taking any other functional norm, for instance the  $L^p$  norm  $\|\cdot\|_p$ .
- In particular, the  $L^2$  norm  $\|\cdot\|_2$  allows use to use Von Neumann stability analysis based on the Fourier transform. This analysis is very convenient and useful for linear problems, but its generalization to nonlinear ones is not easy.

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 |u_i^{n+1}| &= \left| u_i^n \left(1 - a \frac{\Delta t}{\Delta x}\right) + a \frac{\Delta t}{\Delta x} u_{i-1}^n \right| \\
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$$|a| \frac{\Delta t}{\Delta x} \leq 1,$$

is satisfied, we obtain

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- Notice that stability is achieved by **upwinding** and by restricting the time step, so that the **numerical dependency domain** contains the analytical one.
- For instance, in the case of the forward difference method, the value of  $u_i^n$  only depends on the initial data in the nodes belonging to the interval

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- On the other hand, the value of the exact solution only depends on the initial data in  $x_i - at_n$ .
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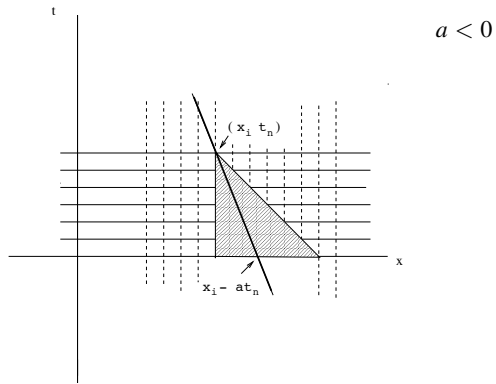
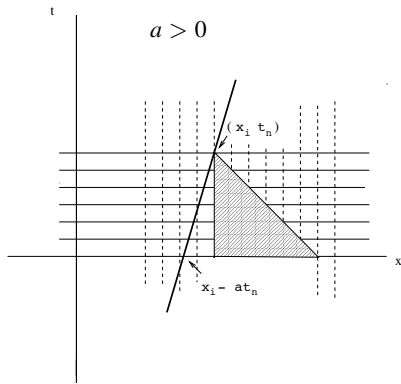
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## Stability



## The upwind method: viscous form

- Using the notation

$$a^+ = \max\{a, 0\}, \quad a^- = \min\{a, 0\}$$

the backward and forward method may be combined in a unique expression:

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which is the so-called **CIR** (after Courant, Isaacson y Rees) or **Upwind method**.

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- It can be interpreted as the centered method plus the discretization of a second order space derivative:

$$\frac{q}{2} \frac{\Delta x}{\Delta t} \Delta x \frac{\partial^2 u}{\partial x^2}.$$

This is the so-called **numerical viscosity**.

- The numerical viscosity stabilizes the centered scheme if the parameter  $q$  is adequately chosen: it can be shown that the numerical method is  $L^2$ -stable if and only if  $q$  is such that:

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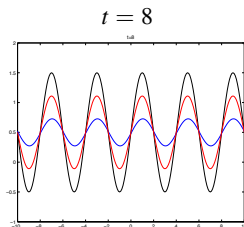
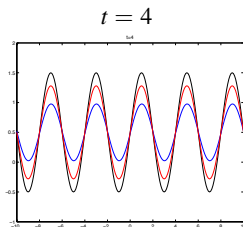
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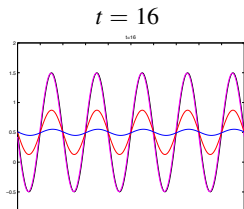
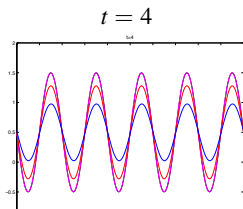
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## Comparison of numerical methods for smooth solutions

First order methods:  $u_0$ ,  $u^n$ -LxF,  $u^n$ -UpW.

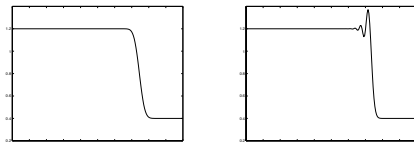


Second order methods: LxW.



## Comparison of numerical methods for discontinuous solutions

- A convergent numerical method can produce wrong numerical approximation of discontinuous solutions.



**Figure:** Numerical solution obtained with the upwind method (left) and the Lax-Wendroff method (right) for the linear equation with  $a = 1$ ,  $\Delta x = 0.01$ ,  $\Delta t = 0.005$

## Nonlinear problems

- Let us consider again the general problem:

$$\frac{\partial u}{\partial t} + \frac{\partial}{\partial x} f(u) = 0, \quad x \in \mathbb{R}, t \geq 0;$$

- If the equation is written in the form:

$$\frac{\partial u}{\partial t} + a(u) \frac{\partial u}{\partial x} = 0,$$

where  $a(u) = f'(u)$ , it is clear that  $a(u)$  locally plays the same role of  $a$  for the linear problem.

- In particular, the following extension of the upwind method is natural:

$$u_i^{n+1} = u_i^n - \frac{\Delta t}{\Delta x} \{a(u_i^n)^+ (u_i^n - u_{i-1}^n) + a(u_i^n)^- (u_{i+1}^n - u_i^n)\}.$$

- This method is consistent, conditionally stable, and of order (1,1). Therefore, it is convergent for smooth solutions. Nevertheless, in the presence of discontinuities the numerical solutions may converge to functions which are not weak solutions of the problem.

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## Finite volume methods

- We define the **cells** or **finite volumes**  $\{I_i\}$  como

$$I_i = [x_{i-1/2}, x_{i+1/2}]$$

where  $x_{i-1/2} = x_i - \Delta x/2$  y  $x_{i+1/2} = x_i + \Delta x/2$ .

- Weak solutions satisfy:

$$\frac{1}{\Delta x} \int_{I_i} u(x, t_{n+1}) dx = \frac{1}{\Delta x} \int_{I_i} u(x, t_n) dx + \frac{\Delta t}{\Delta x} \left\{ \frac{1}{\Delta t} \int_{t_n}^{t_{n+1}} f(u(x_{i-1/2}, t)) dt - \frac{1}{\Delta t} \int_{t_n}^{t_{n+1}} f(u(x_{i+1/2}, t)) dt \right\}.$$

- $u_i^n$  will be interpreted as an approximation of the **average** of  $u$  at the cell  $I_i$  at the instant  $t_n$ :

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- It is natural to consider numerical methods of the form:

$$u_i^{n+1} = u_i^n + \frac{\Delta t}{\Delta x} \{F_{i-1/2}^n - F_{i+1/2}^n\}, \quad (22)$$

where:

$$F_{i+1/2}^n = F(u_{i-q}^n, u_{i-q+1}^n, \dots, u_{i+p-1}^n, u_{i+p}^n).$$

is a consistent approximation of the averaged flow through the intercell  $x_{i+1/2}$  between the times  $t_n$  and  $t_{n+1}$ :

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- If the numerical flux  $F$  is a Lipschitz-continuous and satisfies:

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the numerical method is said to be **consistent**.

- The **Lax-Wendroff's Theorem** ensures that, if the numerical approximation provided by a consistent conservative method converge, the limit is a weak solution of the equation.
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## Finite volume methods

- In the case of a linear equation, all the finite difference methods previously introduced can be interpreted as conservative methods. In particular the centered, upwind, Lax-Wendroff, and Lax-Friedrichs methods are equivalent respectively to the conservative methods corresponding to the numerical fluxes:

$$F^{cen}(u, v) = a \frac{u + v}{2},$$

$$F^{cir}(u, v) = a \frac{u + v}{2} - \frac{|a|}{2}(u - v),$$

$$F^{LW}(u, v) = a \frac{u + v}{2} - \frac{a^2 \Delta t}{2\Delta x}(u - v),$$

$$F^{LF}(u, v) = a \frac{u + v}{2} - \frac{\Delta x}{2\Delta t}(u - v).$$



## Finite volume methods

- It is easy to generalize these numerical fluxes to the general case:

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 F^{cen}(u, v) &= \frac{f(u) + f(v)}{2}, \\
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- A basic requirement is the **linear stability**: if the conservative method is applied to the linear equation, it should be at least conditionally stable. For instance, the generalization of the centered method is not linearly stable and thus it is useless.

## Finite volume methods

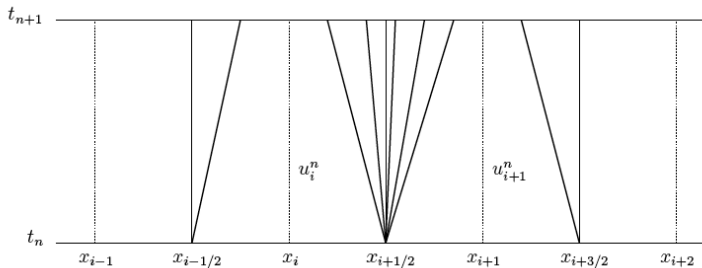
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## Godunov method

- Strategy of the method:



# Godunov method

- Once the approximations at time  $t = t_n$  have been computed, the Cauchy problem

$$\begin{cases} \frac{\partial \tilde{u}}{\partial t} + \frac{\partial}{\partial x} f(\tilde{u}) = 0, & x \in \mathbb{R}, t > t_n, \\ \tilde{u}(x, t_n) = u^n(x), & x \in \mathbb{R}, \end{cases} \quad (23)$$

is considered, where  $u^n$  is again the piecewise constant taking value  $u_i^n$  at the cell  $I_i$ .

- The approximations at time  $t_{n+1}$  are computed then by averaging at every cell the solution of this Cauchy problem:

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- The solution of this Cauchy problem can be computed in terms of the solutions of a family of Riemann problems:

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- If  $f$  is strictly convex or concave, the self-similar solutions of the Riemann problem

$$v(x, t) = V\left(\frac{x}{t}; u_i^n, u_{i+1}^n\right)$$

can be easily computed.

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- If the CFL1/2 condition

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- **Exercise 5:** Apply Godunov method to the linear conservation law

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## Approximate Riemann solvers

- If the computation of the solutions of the Riemann problems is difficult or very costly (this may be the case when  $f$  is not strictly concave or convex, and for systems), a possible strategy is to use an **approximate Riemann solver**: the idea is to use an approximation of the solution of the Riemann problems associated to every inter-cell.

$$\tilde{u}(x, t) = \tilde{V} \left( \frac{x - x_{i+1/2}}{t - t_n}; u_i^n, u_{i+1}^n \right), \quad x \in [x_i, x_{i+1}], \quad t \in (t_n, t_{n+1}].$$

- In the case of the **linear Riemann solvers**,  $\tilde{V}(x/t; u, v)$  is the self-similar solution of a linear Riemann problem:

$$\begin{cases} \frac{\partial \tilde{w}}{\partial t} + a(u, v) \frac{\partial \tilde{w}}{\partial x} = 0, & x \in \mathbb{R}, \quad t > 0, \\ \tilde{w}(x, 0) = \begin{cases} v_L & \text{if } x < 0; \\ v_R & \text{if } x > 0; \end{cases} \end{cases}$$

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## Linear approximate Riemann solvers

- Reasoning as in the case of the Godunov method, if the CFL condition

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is satisfied, where

$$a_{i+1/2}^n = a(u_i^n, u_{i+1}^n),$$

it is possible to rewrite the method as follows:

$$\begin{aligned} u_i^{n+1} &= \frac{1}{\Delta x} \left( \int_{x_{i-1/2}}^{x_i} \tilde{v}(x, t_{n+1}) dx + \int_{x_i}^{x_{i+1/2}} \tilde{v}(x, t_{n+1}) dx \right) \\ &= \frac{1}{\Delta x} \left( \int_{x_{i-1/2}}^{x_i} \tilde{v}(x, t_n) dx + \int_{x_i}^{x_{i+1/2}} \tilde{v}(x, t_n) dx \right. \\ &\quad \left. + \int_{t_n}^{t_{n+1}} a_{i-1/2}^n (\tilde{V}(0; u_{i-1}^n, u_i^n) - u_i^n) dt - \int_{t_n}^{t_{n+1}} a_{i+1/2}^n (u_i^n - \tilde{V}(0; u_i^n, u_{i+1}^n)) dt \right) \\ &= u_i^n + \frac{\Delta t}{\Delta x} (a_{i-1/2}^n (\tilde{V}(0; u_{i-1}^n, u_i^n) - u_i^n) + a_{i+1/2}^n (u_i^n - \tilde{V}(0; u_i^n, u_{i+1}^n))), \end{aligned}$$

## Linear approximate Riemann solvers

- In the previous equality, it has been used that  $\tilde{V}(x/t; u_i^n, u_{i+1}^n)$  is a solution of

$$\tilde{v}_t + a_{i+1/2}^n \tilde{v}_x = 0$$

in  $[x_i, x_{i+1/2}] \times [t_n, t_{n+1}]$  and a solution of

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- Now observe that, if  $a_{i+1/2}^n > 0$  then:

$$a_{i+1/2}^n (u_i^n - \tilde{V}(0; u_i^n, u_{i+1}^n)) = a_{i+1/2}^n (u_i^n - u_{i+1}^n) = 0,$$

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- Therefore:

$$a_{i+1/2}^n (u_i^n - \tilde{V}(0; u_i^n, u_{i+1}^n)) = a_{i+1/2}^{n,-} (u_i^n - u_{i+1}^n).$$

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## Roe method

- Therefore, a numerical method base on a linear approximate Riemann solver can be written in the form:

$$u_i^{n+1} = u_i^n - \frac{\Delta t}{\Delta x} (a_{i-1/2}^{n,+}(u_i^n - u_{i-1}^n) + a_{i+1/2}^{n,-}(u_{i+1}^n - u_i^n)).$$

- This numerical method can be interpreted as a conservative method if the linearization  $a(u, v)$  satisfies the **Roe property**:

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- Using the equalities:

$$a^+ = \frac{a + |a|}{2}, \quad a^- = \frac{a - |a|}{2},$$

we obtain:

$$\begin{aligned} u_i^{n+1} &= u_i^n - \frac{\Delta t}{\Delta x} (a_{i-1/2}^{n,+} (u_i^n - u_{i-1}^n) + a_{i+1/2}^{n,-} (u_{i+1}^n - u_i^n)) \\ &= u_i^n - \frac{\Delta t}{2\Delta x} (a_{i-1/2}^n (u_i^n - u_{i-1}^n) + a_{i+1/2}^n (u_{i+1}^n - u_i^n)) \\ &\quad + |a_{i-1/2}^n| (u_i^n - u_{i-1}^n) - |a_{i+1/2}^n| (u_{i+1}^n - u_i^n) \\ &= u_i^n - \frac{\Delta t}{2\Delta x} (f(u_i^n) - f(u_{i-1}^n) + f(u_{i+1}^n) - f(u_i^n)) \\ &\quad + |a_{i-1/2}^n| (u_i^n - u_{i-1}^n) - |a_{i+1/2}^n| (u_{i+1}^n - u_i^n) \\ &= u_i^n + \frac{\Delta t}{\Delta x} (F^{Roe}(u_{i-1}^n, u_i^n) - F^{Roe}(u_i^n, u_{i+1}^n)) \end{aligned}$$

## Roe property

- where the numerical flux is given by:

$$F^{Roe}(u, v) = \frac{f(u) + f(v)}{2} - \frac{1}{2}|a(u, v)|(v - u).$$

- For scalar problems, the Roe property and the continuity determine the linearization  $a(u, v)$ :

$$a(u, v) = \begin{cases} \frac{f(u) - f(v)}{u - v} & \text{if } u \neq v; \\ f'(u) & \text{if } u = v. \end{cases}$$

- Roe method has very good shock-capturing properties: if the solution of the Riemann problem associated to the states  $u_L, u_R$  consists of a shock linking the states, then the solution of the approximate linear Riemann problem is the same.
- In particular, stationary shock waves are exactly captured.

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- Nevertheless, Roe method has an important drawback: the numerical solutions can converge to a weak solution which is not the entropy solution of the system.
- This drawback can be overcome by using an **entropy fix technique** that prevents the numerical viscosity to vanish when one of the eigenvalues vanishes. For instance, the entropy fix technique proposed in [Harten, Hyman 1983](#) consists in modifying the numerical flux as follows:

$$\tilde{F}_\varepsilon^{Roe}(u, v) = \frac{f(u) + f(v)}{2} - \frac{1}{2} |a(u, v)|_\varepsilon (u - v),$$

where:

$$|a|_\varepsilon = |a| + 0.5 \left\{ \left( 1 + \operatorname{sgn}(\varepsilon - |a|) \right) \left( \frac{a^2 + \varepsilon^2}{2\varepsilon} - |a| \right) \right\},$$

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# Roe method

- **Exercise 7:** Apply Roe method to the linear conservation law

$$u_t + au_x = 0,$$

and check that it coincides with the upwind method.

- **Exercise 8:** Write explicitly the numerical flux of Roe method applied to Burgers equation:

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## Stability and convergence

- Lax-Wendroff's theorem does not ensure the convergence of conservative methods: it only characterizes the possible limits of the numerical solutions.
- In order to ensure the convergence, an adequate stability condition has to be required. caso de los problemas lineales.
- The right concept of stability for nonlinear problems is related to the **total variation**: let  $u^n$  be the piecewise constant function taking value  $u_i^n$  in  $I_i$ . The total variation of  $u^n$  is the quantity:

$$TV(u^n) = \sum_i |u_{i+1}^n - u_i^n|. \quad (24)$$

- A numerical method is said to be **TV-stable** if, for any initial condition  $u_0$  and  $T > 0$ , there exist two positive constant  $\Delta t_0, K$  such that:

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- It can be shown that, if a consistent numerical method is *TV*-stable, the numerical solutions converge to weak solutions of the problem in some sense.

- A numerical method is said to be **TVD** (*Total Variation Diminishing*) if, for every  $u_0$ ,  $\Delta t$ ,  $\Delta x$ , one has:

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## Numerical methods for systems of conservation laws

- Let us consider again a system of conservation laws:

$$\mathbf{u}_t + \mathbf{u}(\mathbf{u})_x = 0.$$

- A conservative numerical method can be written as follows:

$$\mathbf{u}_i^{n+1} = \mathbf{u}_i^n + \frac{\Delta t}{\Delta x} \{ \mathbf{F}_{i-1/2}^n - \mathbf{F}_{i+1/2}^n \},$$

where

$$\mathbf{F}_{i+1/2}^n = \mathbf{F}(\mathbf{u}_i^n, \mathbf{u}_{i+1}^n)$$

is the numerical flux.

- Most of the numerical fluxes introduced for scalar conservation laws can be easily extended to systems. This is the case of the Lax-Wendroff or the Lax-Friedrichs fluxes:

$$\begin{aligned} \mathbf{F}^{LW}(\mathbf{u}, \mathbf{v}) &= \frac{\mathbf{f}(\mathbf{u}) + \mathbf{f}(\mathbf{v})}{2} - \frac{\Delta t}{2\Delta x} \mathcal{A}((\mathbf{u} + \mathbf{v})/2)^2 (\mathbf{u} - \mathbf{v}), \\ \mathbf{F}^{LF}(\mathbf{u}, \mathbf{v}) &= \frac{\mathbf{f}(\mathbf{u}) + \mathbf{f}(\mathbf{v})}{2} - \frac{\Delta x}{2\Delta t} (\mathbf{u} - \mathbf{v}). \end{aligned}$$

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represents the self-similar solution of the Riemann problem with initial data  $\mathbf{u}$  and  $\mathbf{v}$ . But it may be difficult to implement. . .

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## Numerical methods for systems of conservation laws

- A Roe linearization is available for the homogeneous shallow water system. Given two states:

$$\mathbf{u}_i = \begin{pmatrix} h_i \\ q_i \end{pmatrix}, \quad i = 1, 2,$$

let us define:

$$\mathcal{A}(\mathbf{u}_1, \mathbf{u}_2) = \begin{bmatrix} 0 & 1 \\ -\bar{u}^2 + g\bar{h} & 2\bar{u} \end{bmatrix} \quad (25)$$

where

$$\bar{u} = \frac{\sqrt{h_1 u_1} + \sqrt{h_2 u_2}}{\sqrt{h_1} + \sqrt{h_2}}, \quad \bar{h} = \frac{h_1 + h_2}{2}, \quad (26)$$

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## Numerical methods for balance laws

- In general, a numerical scheme is said to be well-balanced if it captures correctly the smooth stationary solutions of the system, or at least a family of them.
- In shallow water systems, the numerical schemes are usually required to preserve at least the solutions corresponding to water at rest: this is the so-called the C-property (Bermúdez & Vázquez, 1994).
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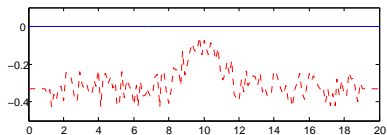
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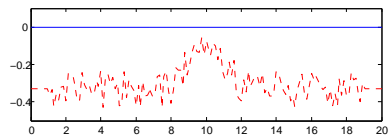
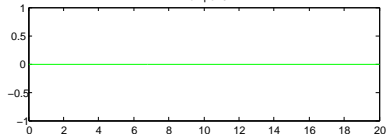
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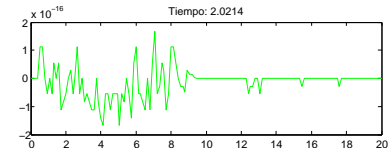
Well-balanced method: **bottom**, **water surface**, **discharge**.



Tiempo: 0

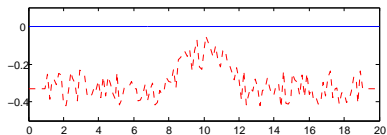


Tiempo: 2.0214

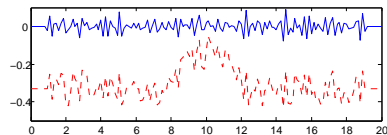
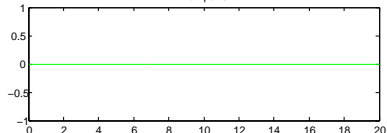


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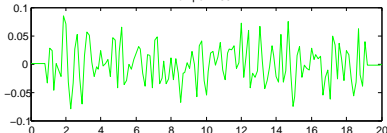
A method which is not well-balanced method: **bottom**, **water surface**, **discharge**.



Tiempo: 0



Tiempo: 2.0012





## Water at rest solutions

- We consider the complete shallow water model:

$$\mathbf{u}_t + \mathbf{f}(\mathbf{u})_x = \mathbf{S}(\mathbf{u}) \frac{dH}{dx},$$

$$\mathbf{u} = \begin{pmatrix} h \\ q \end{pmatrix}, \quad \mathbf{f}(\mathbf{u}) = \begin{pmatrix} q \\ q^2/h + gh^2/2 \end{pmatrix}.$$

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- It can be easily shown that, given any constant  $C$ ,

$$\mathbf{u}(x) = \begin{pmatrix} h(x) \\ q(x) \end{pmatrix} = \begin{pmatrix} C + H(x) \\ 0 \end{pmatrix}, \quad \forall x,$$

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# Numerical methods

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## Well-balanced property

- The numerical method is said to satisfy the  $C$ -property if it preserves water at rest solutions.
- More specifically, given any water at rest solution

$$h = H + C, \quad q = 0,$$

if we consider the initial condition

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$$S^\pm(\mathbf{u}, \mathbf{v}) = \frac{1}{2} \mathbf{S} \left( \frac{1}{2}(\mathbf{u} + \mathbf{v}) \right),$$

that is, a centered discretization of the source term. Nevertheless, with this choice the numerical method is not well-balanced.

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## Well-balanced property

- Let us consider the matrices:

$$\mathcal{P}^{\pm}(\mathbf{u}, \mathbf{v}) = \frac{1}{2} \mathcal{K}(\mathbf{u}, \mathbf{v}) \cdot (I \pm \text{sgn}(\Lambda(\mathbf{u}, \mathbf{v}))) \cdot \mathcal{K}(\mathbf{u}, \mathbf{v})^{-1},$$

where  $I$  is the identity matrix and  $\text{sgn}(\Lambda(\mathbf{u}, \mathbf{v}))$  is the diagonal matrix whose diagonal coefficients are the sign of the eigenvalues.

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- Exercise 11:** Prove the equalities:

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## Upwind discretization of the source term

- We propose the definitions:

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## Upwind discretization of the source term

- Finally, using the Roe property and the equality (27), we obtain:

$$\begin{aligned}
 \mathbf{u}_i^1 &= \mathbf{u}_i^0 - \frac{\Delta t}{\Delta x} \left( \right. \\
 &\quad \mathcal{P}_{i-1/2}^{0,+} \left( \mathbf{f}(\mathbf{u}_i^0) - \mathbf{f}(\mathbf{u}_{i-1}^0) + S_{i-1/2}(H_i - H_{i-1}) \right) + \\
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 &= \mathbf{u}_i^0,
 \end{aligned}$$

as we wanted to prove.

## Upwind discretization of the source term

- It is possible to derive well-balanced numerical methods that use any numerical flux: [Audusse, Bouchut, Bristeau, Klein, Perthame 2004](#), [Castro, Pardo, CP, Toro, 2011...](#)
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## Extension to high order

- Let us consider a **reconstruction operator** that, given a sequence of cell values  $\{\mathbf{u}_i\}$ , provide a smooth function at every cell

$$P_i(x; \mathbf{u}_{i-l}, \dots, \mathbf{u}_{i+r})$$

that depend on the values at the neighbor cells  $\{I_k\}_{k=i-l}^{i+r}$  that constitute the *stencil*, in such a way that:

- It is **conservative**, i.e.

$$\mathbf{u}_i = \frac{1}{\Delta x} \int_{x_{i-1/2}}^{x_{i+1/2}} P_i(x; \mathbf{u}_{i-l}, \dots, \mathbf{u}_{i+r}) dx.$$

- If the cell values are the averages of a smooth function  $\mathbf{u}(x)$  then

$$\mathbf{u}(x) = P_i(x; \mathbf{u}_{i-l}, \dots, \mathbf{u}_{i+r}) + O(\Delta x^p), \quad \forall x \in [x_{i-1/2}, x_{i+1/2}].$$

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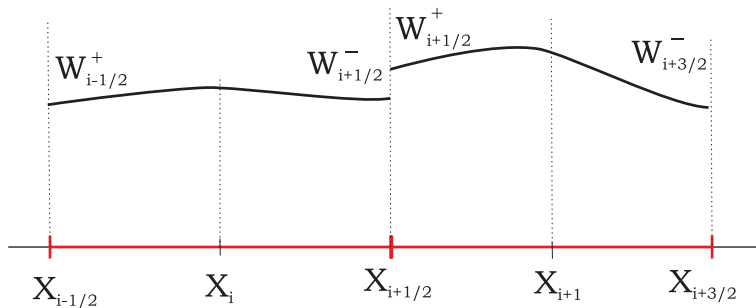
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- Let us suppose that the depth function  $H$  is continuous. We consider the semidiscret numerical method:

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- The numerical method is well-balanced for water at rest if the reconstruction operator is exact for constant functions and the reconstruction is computed as follows:
  - First, the reconstruction operator is applied to the variables  $q$  and  $\eta = h - H$  to obtain  $p'_{q,i}, p'_{\eta,i}$ .
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- The semidiscrete scheme is fully discretized by using an adequate time-stepping, for instance TVD-RD methods.
- The integral terms can be computed with a quadrature formula with an order of accuracy greater or equal than  $p$ .
- Examples of reconstruction operators: MUSCL (Van Leer 1979), PPM (Colella, Woodward 1984), ENO (Harten, Engquist, Osher & Chakravarthy, 1987), WENO (Liu, Osher & Chan, 1994, Jiang & Shu, 1996), PHM (Marquina, 1994), etc.
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  - Then, the reconstruction of the variable  $h$  is defined by  $p_{h,i}^t = H + p_{\eta,i}^t$ .
- The semidiscrete scheme is fully discretized by using an adequate time-stepping, for instance TVD-RD methods.
- The integral terms can be computed with a quadrature formula with an order of accuracy greater or equal than  $p$ .
- Examples of reconstruction operators: MUSCL (Van Leer 1979), PPM Colella, Woodward 1984), ENO (Harten, Engquist, Osher & Chakravarthy, 1987), WENO (Liu, Osher & Chan, 1994, Jiang & Shu, 1996), PHM (Marquina, 1994), etc.
- It is possible to obtain high-order methods that preserve every stationary solution and not only water at rest ones: Noelle, Pankratz, Puppo, Natvig, 2006, Noelle, Xing, Shu, 2007, Russo, Khe, 2009, Castro, López, CP, 2013...

## Extension to 2d problems

- The 2d problem can be written in the form:

$$\mathbf{u}_t + \mathbf{f}_1(\mathbf{u})_x + \mathbf{f}_2(\mathbf{u})_y = S_1(\mathbf{u})H_x + S_2(\mathbf{u})H_y,$$

- where

$$\mathbf{u} = \begin{pmatrix} h \\ q_1 \\ q_2 \end{pmatrix}, \quad \mathbf{f}_1(\mathbf{u}) = \begin{pmatrix} q_1 \\ \frac{q_1^2}{h} + \frac{g}{2}h^2 \\ \frac{q_1 q_2}{h} \end{pmatrix}, \quad \mathbf{f}_2(\mathbf{u}) = \begin{pmatrix} q_2 \\ \frac{q_1 q_2}{h} \\ \frac{q_2^2}{h} + \frac{g}{2}h^2 \end{pmatrix},$$

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# Finite volumes

- The domain  $D$  is decomposed into subsets (closed polygons) called cells of finite volumes,  $V_i \subset \mathbb{R}^2$ ;
- $N_i \in \mathbb{R}^2$  is the centre of  $V_i$ ;
- $\mathcal{N}_i$  is the set of indexes  $j$  such that  $V_j$  is a neighbor of  $V_i$ ;
- $E_{ij}$  is the common edge to two neighbor cells  $V_i$  and  $V_j$ , and  $|E_{ij}|$  represents its length;
- $\eta_{ij} = (\eta_{ij,1}, \eta_{ij,2})$  is the normal unit vector of the edge  $E_{ij}$  pointing towards the cell  $V_j$ ;

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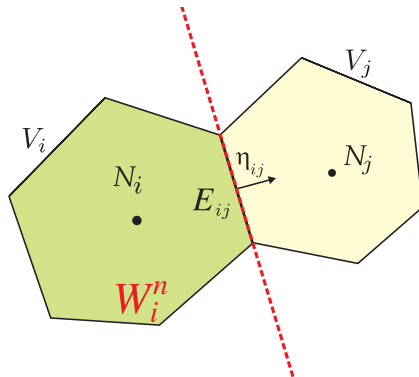
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- The numerical scheme writes as follows:

$$\mathbf{u}_i^{n+1} = \mathbf{u}_i^n - \frac{\Delta t}{|V_i|} \left( \sum_{j \in \mathcal{N}_i} |E_{ij}| (F_{ij} - S_{ij}^- (H_j - H_i)) \right)$$

where

$$F_{ij} = \eta_{ij,1} F_1(\mathbf{u}_i^n, \mathbf{u}_j^n) + \eta_{ij,2} F_2(\mathbf{u}_i^n, \mathbf{u}_j^n);$$

$$S_{ij} = \eta_{ij,1} S_1^-(\mathbf{u}_i^n, \mathbf{u}_j^n) + \eta_{ij,2} S_2^-(\mathbf{u}_i^n, \mathbf{u}_j^n).$$

- For the extension of high-order methods to 2d problems, see [Castro, Fernández-Nieto, Ferreiro, García, CP, 2009](#).

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